Graduate Physics Seminar:

Bora wind in Slovenia

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Outline

1 Introduction

2 Downslope winds

3 Wind field measuring techniques

4 Case study of Bora wind in Vipava valley

What is Bora wind?

- North-eastern or East-north-eastern wind
- Caused by the flow over orographic barrier
- Cold lee-side/downslope wind
- Strong wind gusts over 50 m/s



Source: http://www.treehouse-maps.com/

Wind measurements

Case study

Conclusions

Effects of strong Bora wind







Source: www.tic-ajdovscina.si, www.siol.net, www.24ur.com

Classification of Bora wind events

- Historic classifications (tree deformation, local measurements,...)
- Synoptic classification: Anticyclonic (Light) Bora:
 - planetary through over West Europe
 - high pressure over Central Europe

Cyclonic (Dark) Bora:

- cyclone in the North Adriatic
- anticyclone over Central Europe





Source: Menezes, A., Tabeaud, M., 2000: Variations in Bora weather type in the north Adriatic See, 1866 - 1998. Weather, vol. 55, 2000, 452-458. University Press.

Downslope winds



- 1 Stably stratified cold air is forced to rise over a topographic barrier
- 2 The air on the lee-side oscillates and forms the so-called mountain waves – inner gravity waves



- **3** Consequences:
 - a drag to the upper level of atmosphere
 - possible clear-air turbulence (CAT)
 - strong surface winds at the lee side of the mountains

Airmass motion - basic set of equations

- Continuity equation: $\frac{D\rho}{Dt} + \rho \nabla \vec{u} = 0 \rightarrow$ no airmass sources or sinks
- Ideal gas equation: $p = \rho RT = \frac{1}{2}RT \rightarrow$ we assume the air to be ideal gas
- Equations of motion:

 $\frac{Du}{Dt} = -\frac{1}{\rho}\frac{\partial p}{\partial x} + fv + F_{rx}$ $\frac{Dw}{Dt} = -\frac{1}{2}\frac{\partial p}{\partial z} - g$

air parcels affected by $\frac{Dv}{Dt} = -\frac{1}{
ho}\frac{\partial p}{\partial y} - fu + F_{ry} \rightarrow$ basic Newtonian mechanics – Newton's 2nd Law

Thermodynamic equation:

 $\frac{1}{\theta_0} \left(\frac{\partial \theta}{\partial t} + u \frac{\partial \theta}{\partial x} + v \frac{\partial \theta}{\partial u} \right) + w \frac{d \ln \theta_0}{dz} = \frac{J}{c_p T} \rightarrow \text{ describes energy}$ sources and sinks

Solutions of basic equations

• Basic equations form a system of non-linear partial differential equations

- Possible approaches to solutions of the system of partial equations:
 - Numerical solution of equations \rightarrow weather forecast
 - Linearization of equations \rightarrow analytical solution, yielding the properties of gravity waves with different wavelengths

Numerical solutions - Weather prediction

• Spectral method – finite series of orthogonal functions to represent the spatial variations (ALADIN)





• Finite difference method – simple discretization, derivatives of the variable are approximated by finite differences

Analytical solution for Bora wind case - Wave forms

- Linearization of partial differential equations with perturbation method
- Solutions can be written in wave form
- Special case inertia-gravity waves over sinusoidal topography:

$$\left(\frac{\partial^2 w^{'}}{\partial x^2} + \frac{\partial^2 w^{'}}{\partial z^2}\right) + \frac{N^2}{\bar{u}^2}w^{'} = 0 \rightarrow \text{ harmonic oscilator}$$



Source: Holton, J. R., 2004: An Introduction to Dynamic Meteorology, Fourth Edition. Oxford: Elsevier Academic Press.

Mountain waves - Downslope windstorm

- Reflection of vertically propagating linear gravity waves in lower layer from the upper layer
- Nonlinear processes have to be taken into account
- Froude number $({\rm Fr}^2=\bar{u}^2/c^2)$ the ratio between mean-flow speed and wave speed



Source: Holton, J. R., 2004: An Introduction to Dynamic Meteorology, Fourth Edition. Oxford: Elsevier Academic Press.

• Critical flow at the crest of the mountain, supercritical flow down the lee side of the mountain, subcritical conditions in the valley \rightarrow downslope windstorm

Wind field measurements

- Changes in wind speed often abrupt:
 - High sampling rate analysis of wind gusts
 - Averaged values of wind speed input for numerical models (initial conditions)
- Measurements of vertical wind profiles, 2D and 3D wind field
 - \rightarrow verification of numerical models



Measurements obtained by Honzak, L., Rakovec, J., Skok, G., Žabkar, R., University of Ljubljana, Faculty of Mathematics and Physics.

Wind measurements

In-situ measurements vs. remote sensing

A point measuremen Meteorological

tower

Measuring techniques:

- In-situ methods:
 - Measurements at or near the ground
 - Sensors measure the properties of atmosphere directly
 - Point measurements
 - Wind speed averaged value over a period of time
 - Gust speed maximum value in the same period

Wind measurements

Conclusions

In-situ measurements vs. remote sensing



Measuring techniques:

- Remote sensing methods:
 - Detection of a signal scattered on atmospheric features (clouds, aerosols, molecules)
 - Passive methods scattering of natural radiation (i. e. sunlight) on objects
 - Active methods system emits the signal and observes the return
 - Volume measurements

In-situ instruments for wind field measurements

- Cup anemometer
- Propeller anemometer
- Hot-wire anemometer
- Wind vane wind direction
- Ultrasonic anemometer 2D or 3D wind vector (measures the time of flight of acoustic signal)



Sources: Cup anemometer - http://www.logotronic.at, Propeller anemometer - http://www.tes.com.tw, Hot-wire anemometer http://www.test-equipment.com.au, Wind vane - http://www.geographylwc.org.uk, Ultrasonic anemometer - http://www.wikipedia.org.

In-situ instruments for wind field measurements

- Radiosondes:
 - vertical profile of 2D wind vector
 - poor temporal resolution two launches per day
 - poor spatial resolution



Sources: Radiosonde - http://radiosondemuseum.org, Radiosonde stations - http://www.eumetnet.eu/.

Remote sensing systems for wind field measurements

- Remote sensing systems:
 - SOund Detecting And Ranging SODAR
 - RAdio Detecting And Ranging RADAR
 - Synthetic Aperture Radar SAR
 - LIght Detecting And Ranging LIDAR



Source: Emeis, S., 2011: Surface-Based Remote Sensing of the Atmospheric Boundary Layer. London: Springer.

Basic lidar configuration



Lidar device:

- Transmitter high energy light source
- Receiver optical system and photomultiplier



Transmitter - receiver configuration:

- Co-axial axis of transmitter and receiver coincide
- Bi-axial axis of transmitter and receiver are separated

Conclusions

Wind lidar measurements



VAD scanning mode

Possible lidar system configurations:

- Doppler lidar
- Fast scanning elastic-backscatter lidar



- Types of scans used to obtain wind field:
 - Time-to-Height Indicator (THI) mode
 - Range Height Indicator (RHI) mode
 - Velocity Azimuth Display (VAD) mode
 - Plan Position Indicator (PPI) mode

Wind measurements

Conclusions

Doppler lidar

- Scattering on molecules \rightarrow frequency shift of the backscattered signal
- Frequency shift → wind speed:
 - Radial component of wind \rightarrow directly from the Doppler shift of the signal
 - 3D wind field \rightarrow scanning at various azimuth angles



- Direction of the largest frequency shift \rightarrow wind direction
- Frequency evaluation:
 - Coherent detection measuring the beat frequency of the signal
 - Direct (incoherent) detection the frequency change determined by transmission change

Wind field measurements - Doppler lidar



An example of measurements Doppler lidar:

- Spatial resolution: 100 m
- Range: from 400 m to 10 km
- Maximum detectable wind speed: 40 m/s

The system is capable of detecting:

- the terrain-induced wind shear
- hydraulic jumps
- vortexes

Wind measurements

Conclusions

Fast scanning elastic-backscatter lidar



- Correlation method → aerosol features identified and followed while drifting with the wind
- Aerosols small and light enough to be assumed they drift with the wind speed
- Measurements strongly dependent on the aerosol concentrations:
 - Low concentrations weak return signal
 - Low clouds, fog or precipitation information about the speed of raindrops instead of wind speed

Source: He, T.-Y., Stanič, S., Gao, F., Bergant, K., Veberič, D., Song, X.-Q., and Dolžan, A, 2012: Tracking of urban aerosols using combined LIDAR-based remote sensing and ground-based measurements. Atmospheric Measuring Techniques, vol. 5, 2012, 891 - 900.

Wind field measurements – fast scanning lidar

An example of measurements of a 2D vector wind fields:

- Scanning speed: 3.75 $^{\circ}/s$ for 90 $^{\circ}$ PPI scan
- Range: 18 km
- Resolution: 15 m



Source: Mayor, S. D., Eloranta, E. W., 2001: Two-Dimensional Vector Wind Fields from Volume Imaging Lidar Data. Journal of Applied Meteorology, vol. 40, August 2001, 1331 - 1346.

Synthetic Aperture Radar

- Coherent sidelooking radar system
- Airborne or spaceborne
- Using the flight path to simulate extremely large antenna
- High-resolution imagery
- Wind field obtained by wind scatterometer model



Source: Measuring scheme – http://www.crisp.nus.edu.sg, Satellite image – Alpers, W., Ivanov, A., Horstmann, J., 2008: Observations of Bora Events over Adriatic Sea and Black Sea by Spaceborne Synthetic Aperture Radar. *Monthly Weather Review*.

Wind field measurements – synthetic aperture radar An example of measurements of wind field with SAR:

- cloud patterns gravity waves, wave-induced rotors
- visible flows through the cols
- wind jets and wakes



Source: Alpers, W., Ivanov, A., Horstmann, J., 2008: Observations of Bora Events over Adriatic Sea and Black Sea by Spaceborne Synthetic Aperture Radar. Monthly Weather Review.

Case study of Bora wind

- Period: 27 January 24 April 2012
- Wind and gust speed measured at 7 locations:
 - 5 locations arranged in two lines perpendicular to the barrier (sampling rate 10 s)
 - 2 meteorlogical stations Otlica and Dolenje (SR 0.5 h)



Conclusions

Vertical profiles through measuring locations



The aim of positioning the sensors in perpedicular lines \rightarrow capturing the downslope acceleration of the air-masses

Orographic features of the area:

- 900 1200 m a.s.l. Trnovski gozd plateau
- \sim 100 m a.s.l. Vipava valley
- ~300 m a.s.l. Karst plateau

Statistics of Bora wind events



- Larger wind speed values for locations in the valley (A02, A03, B02) than on the slope (A01, B01)
- Lowest wind speed at A01
- Peaks of hourly-mean wind speed coincide for all the measuring points

Limitations of measuring system and possible solutions

- Spatial and temporal dynamics of the processes numerical models
- Verification of numerical models:
 - In-situ measurements low spatial and temporal resolution
 - Lidar system spatial resolution of a few meters and temporal resolution of a few seconds
- Requirements for the lidar system:
 - small device mobility
 - scanning in two directions (zenith and azimuth angle)
 - fast scanning
 - fast data acquisition

Future work



A small rapid scanning lidar system:

- Transmitter Nd:YAG pulsed laser (532 nm, 5 kHz repetition rate, energy 3 μJ/pulse)
- Receiver Ritchey-Crétien optic system with two hyperbolic mirrors
- Bi-axial configuration of transmitter and receiver
- Support frame movements in alt-azimuthal direction
- Light return collection photomultiplier
- Fast data acquisition custom made electronics

Conclusions

Retrieval of wind information at high frequency and at different heights:

 Ultrasonic anemometers → better insight into the frequency and eventual periodicity of the wind gusts

- Fast scanning elastic lidar:
 - vertical wind profiles \rightarrow depth of the Bora wind flow
 - 2D wind profiles \rightarrow structure of the Bora wind flow



